

HyPOP: Momentum and Tracers on Separate Vertical ALE Grids

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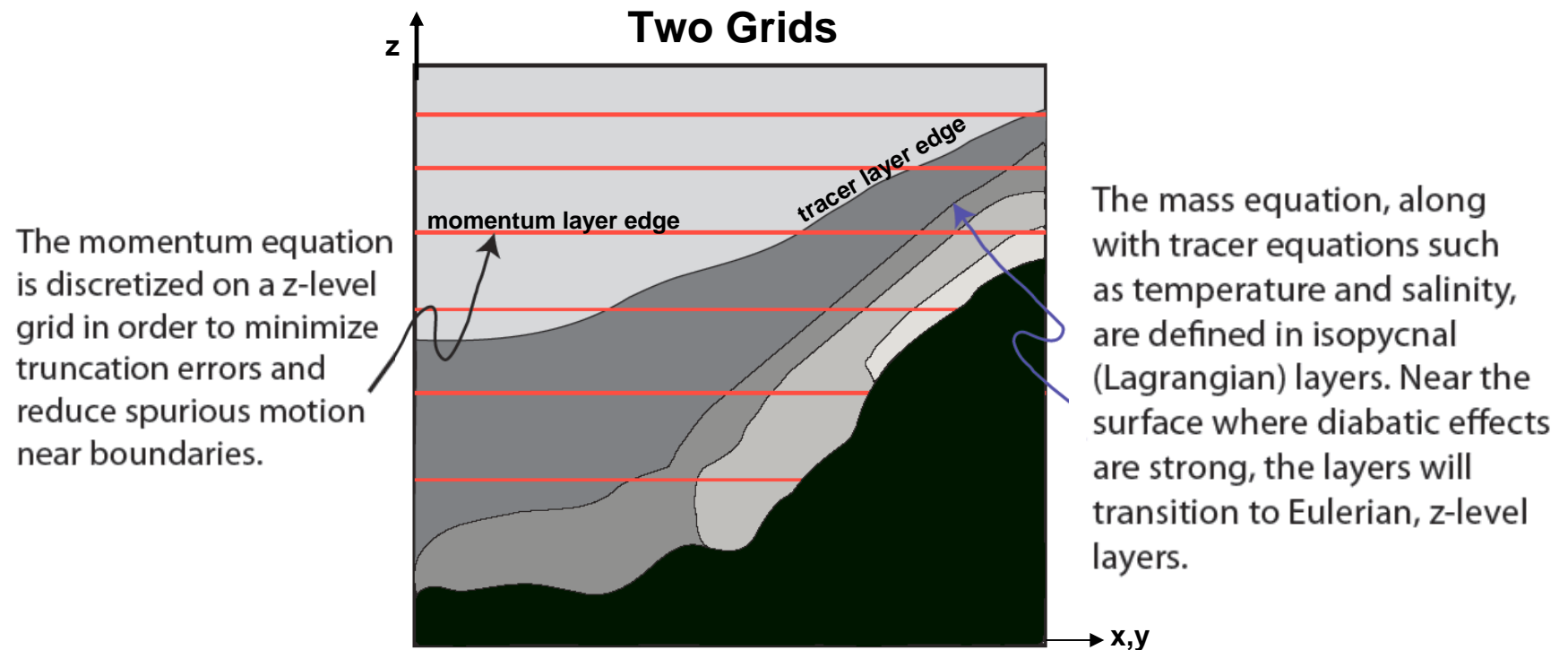


Slide 1

HyPOP: Motivation

- **Isopycnal grids have advantages for tracer transport:**
 - Transport and mixing in the deep ocean follows isopycnal surfaces, so isopycnal-based grids have less diapycnal mixing at depth.
 - Overflow regions better represented in isopycnal grids.
 - Z-grid models, like POP, require corrections to better represent this isopycnal flow and mixing.
- **Ice shelf/ocean interaction better modeled with layer grids**
- **We want to retain the well-tested, z-level POP formulation for the momentum equation.**
- **HyPOP solves momentum equation and tracer equations on different grids, allowing great versatility.**

HyPOP: Momentum Grid and Tracer Grid



HyPOP: Momentum Grid and Tracer Grid

Momentum Grid

conservation of momentum

$$\partial_t \mathbf{u} + \underbrace{\mathbf{u} \cdot \nabla \mathbf{u}}_{\text{advection}} - \underbrace{f \times \mathbf{u}}_{\text{Coriolis}} = \underbrace{-\rho_0^{-1} \nabla p}_{\text{pressure gradient}} + \underbrace{A_M \nabla_h^2 \mathbf{u} + \partial_z \mu \partial_z \mathbf{u}}_{\text{diffusion}}$$

Interpolation Between Grids

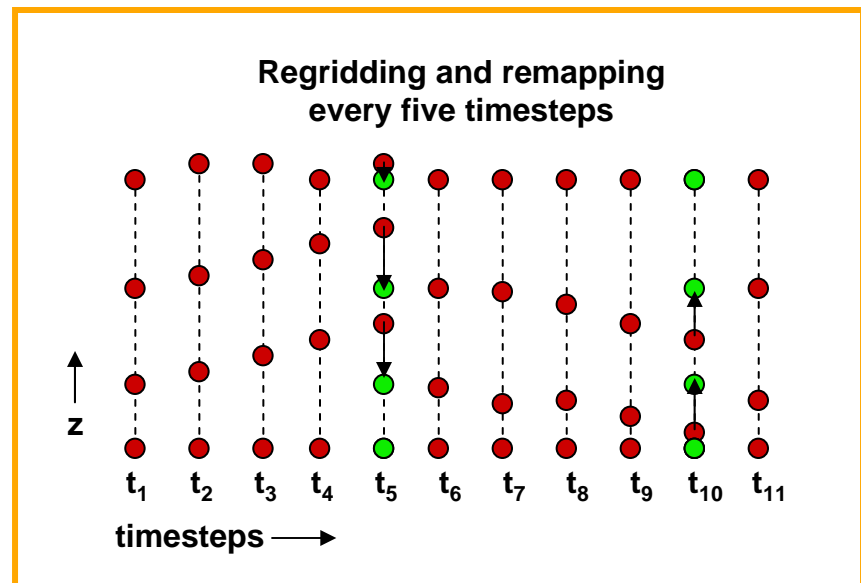
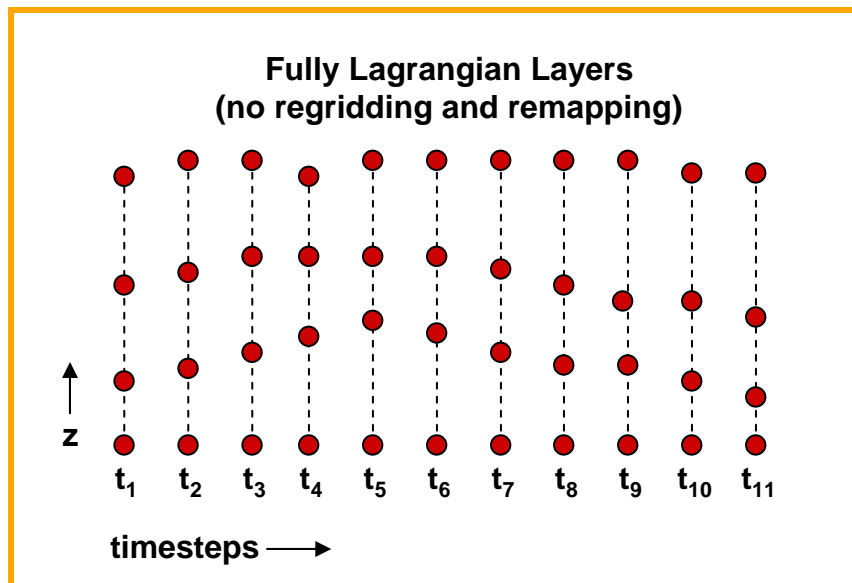
Tracer Grid

conservation of tracers (temperature, salinity, etc)

$$\partial_t \varphi + \underbrace{\mathbf{u} \cdot \nabla \varphi}_{\text{advection}} = \underbrace{A_H \nabla_h^2 \varphi + \partial_z \kappa \partial_z \varphi}_{\text{diffusion}} + \underbrace{Q}_{\text{source/sink}}$$

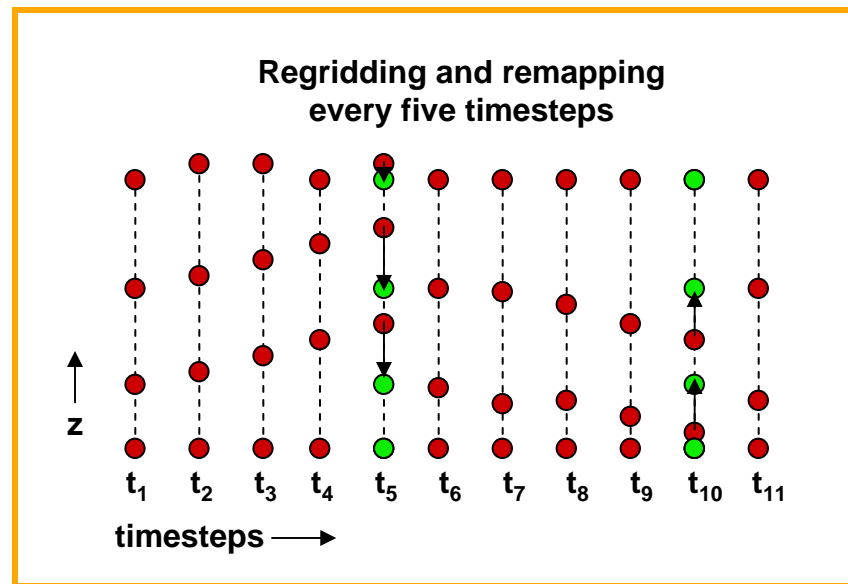
ALE: Arbitrary Lagrangian-Eulerian Coordinates

- Isopycnal coordinates are naturally Lagrangian; isopycnal surfaces move with the fluid.
- ALE algorithm consists of three parts:
 1. Lagrangian steps, where vertical grid moves with the fluid.
 2. Regridding step, where grid is modified to ensure it is smooth and well-spaced.
 3. Remapping step, where variables are conservatively transformed from the old grid to the new.



ALE: Targets for Regridding

- **Regridding and remapping may be done:**
 - after every step
 - after every n steps, or
 - based on grid criteria
- **Grid criteria that initiate regridding and remapping would include:**
 - Layers that deviate too far from the target grid
 - Layers that are too thin or too thick
 - Layer interfaces that are not smooth



HyPOP: Versatility in both Tracer and Momentum Grids

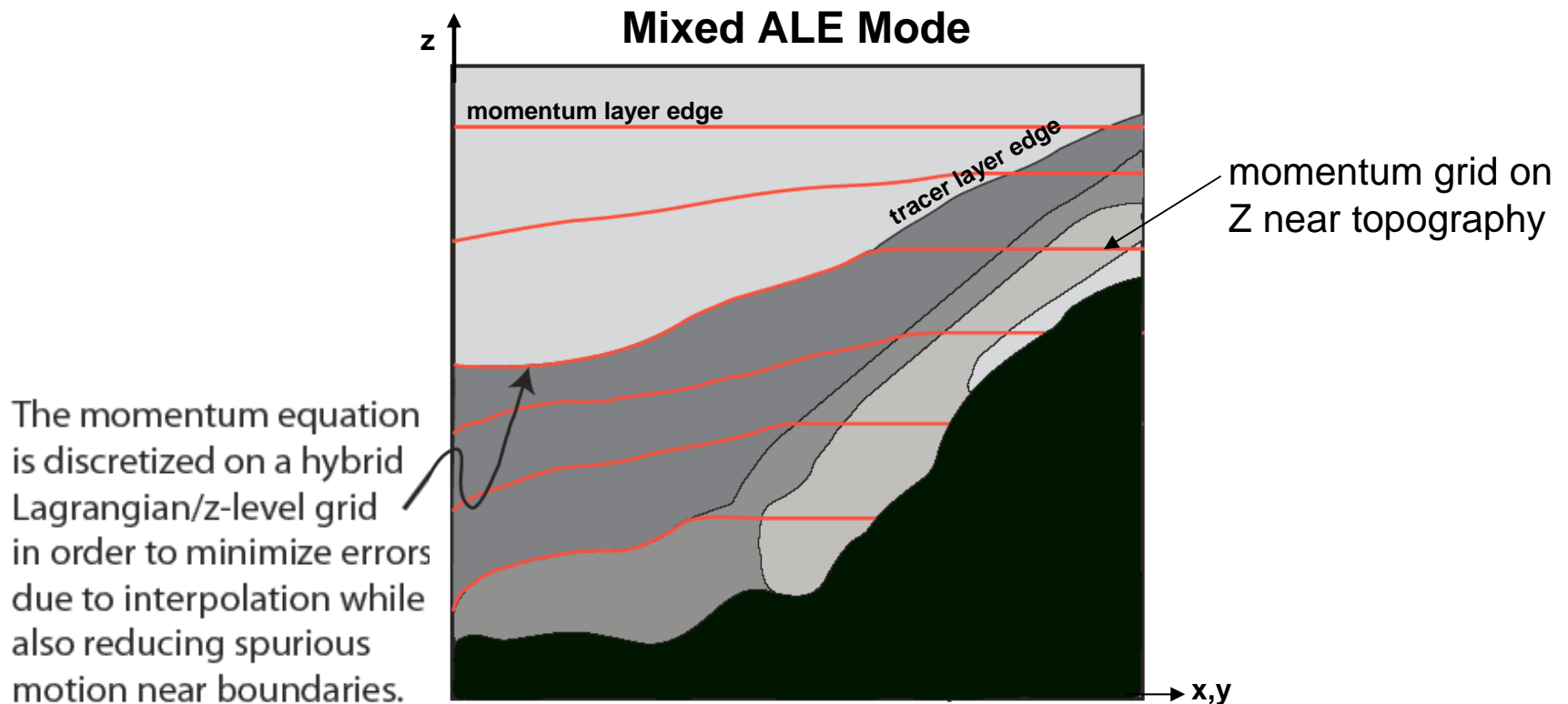
- HyPOP's software infrastructure allows great flexibility in vertical grids.
- The tracer and momentum grids may each be Z-level or ALE.
- Possible configurations include:

Tracer grid: Momentum grid:	Z-level	ALE (isopycnal, Z on top)
Z-level	POP mode	'Standard' HyPOP mode
ALE (isopycnal, Z on top)	NA	hybrid layered model
ALE (Z near topography)	NA	mixed ALE mode

- This framework allows us to make quantitative comparisons to test:
 - Improvements in tracer advection and diffusion on a Z versus isopycnal grid.
 - Additional computation required for interpolation when two grids are used.

HyPOP: Versatility in both Tracer and Momentum Grids

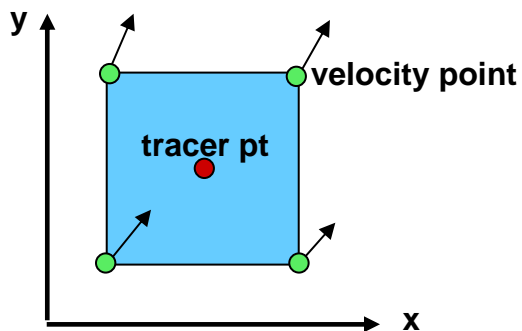
Tracer grid: \ Momentum grid:	Z-level	ALE (isopycnal, Z on top)
	POP mode	'Standard' HyPOP mode
Z-level	POP mode	'Standard' HyPOP mode
ALE (isopycnal, Z on top)	NA	standard layered model
ALE (Z near topography)	NA	mixed ALE mode



Other Details

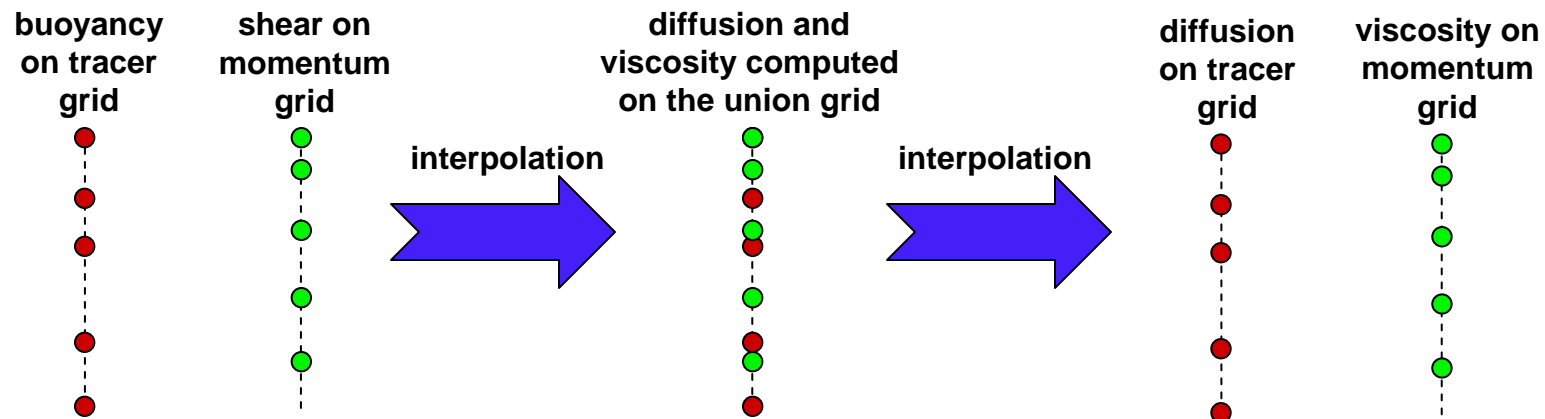
	POP	HyPOP
Horizontal Grid	B-grid	B-grid
Horizontal Advection	centered diff.	incremental remapping centered diff. still an option
Timestepping	leapfrog	leapfrog Two-level schemes to be implemented later.
Barotropic/baroclinic splitting	Split implicit/explicit	Split implicit/explicit

B-grid: Velocities on corners



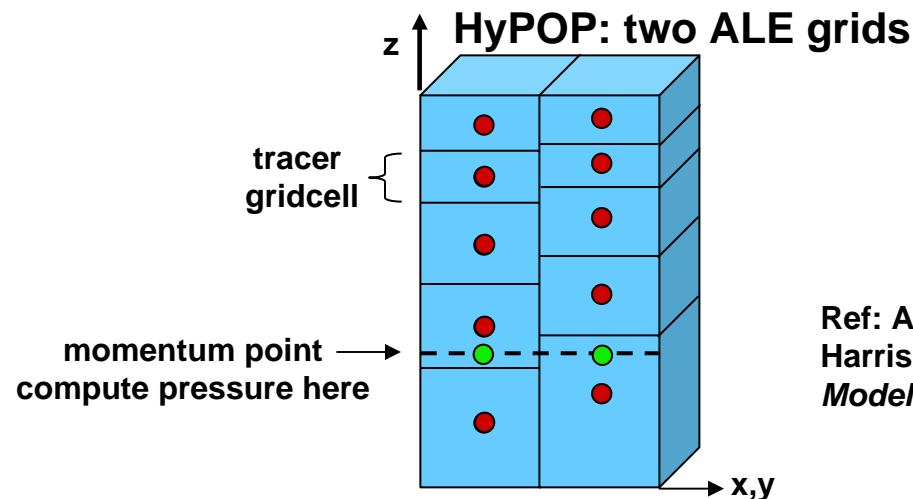
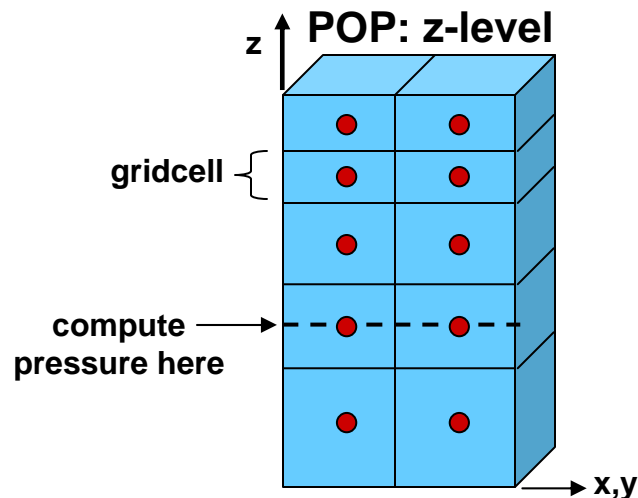
HyPOP: Vertical diffusion conducted on Union Grid

- Vertical diffusion coefficients are needed on the tracer grid.
- Viscosity is needed on the momentum grid.
- Both are computed using parameterizations (KPP, Richardson number) that use:
 - Shear from the momentum grid
 - Buoyancy from the tracer grid
- Parameterizations are implemented on a Union grid.
- This avoids loss of accuracy due to interpolation from tracer to momentum grid, and vice versa.



HyPOP: How do we avoid the thermobaric instability?

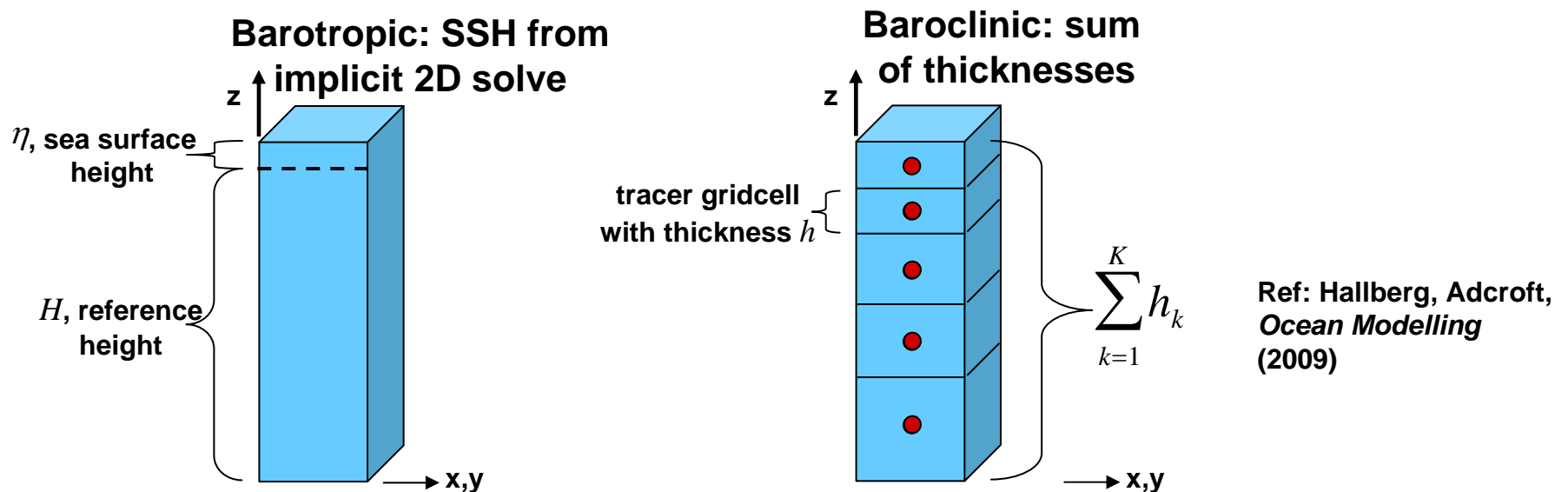
- The density is a function of pressure (depth) as well as T and S .
- POP: computes $p(z) = -\int_0^z \rho g dz'$ as a simple linear sum, $p_K = -g \sum_{k=1}^K \rho_k \Delta z_k$
- This works because neighboring columns have gridpoints at the same depth.
- HyPOP: Like HYCOM, neighboring columns have gridpoints at different depths, so computation of pressure must be more accurate. If not, spurious pressure gradients form.
- We construct a cubic spline of density in each column, and compute p at each momentum level by integration of the cubic spline.



Ref: Adcroft, Hallberg, Harrison, *Ocean Modelling* (2008)

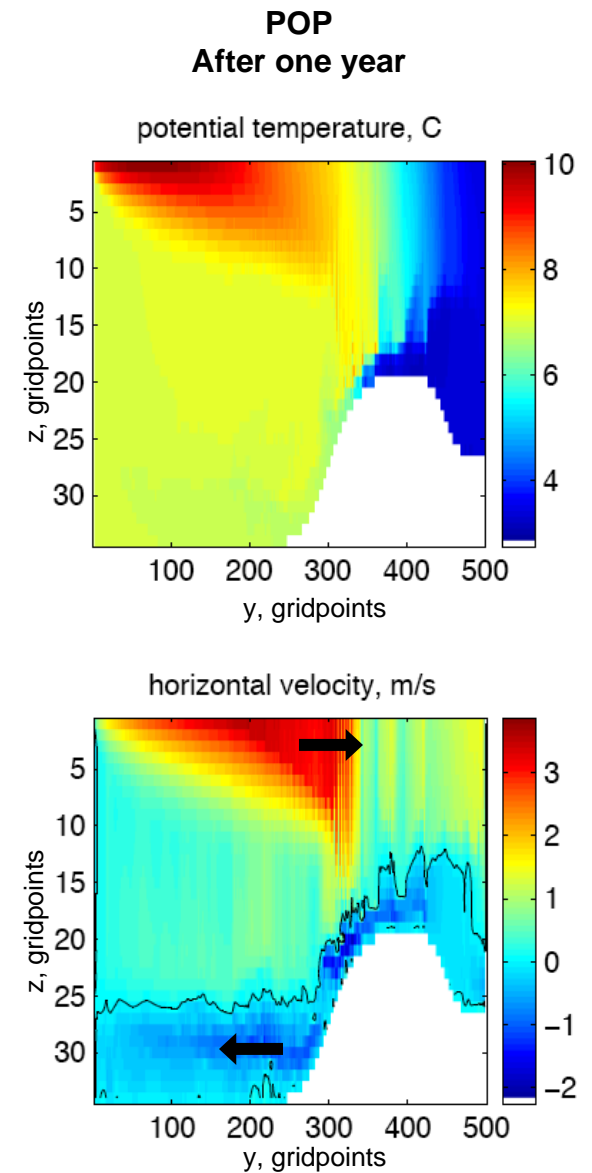
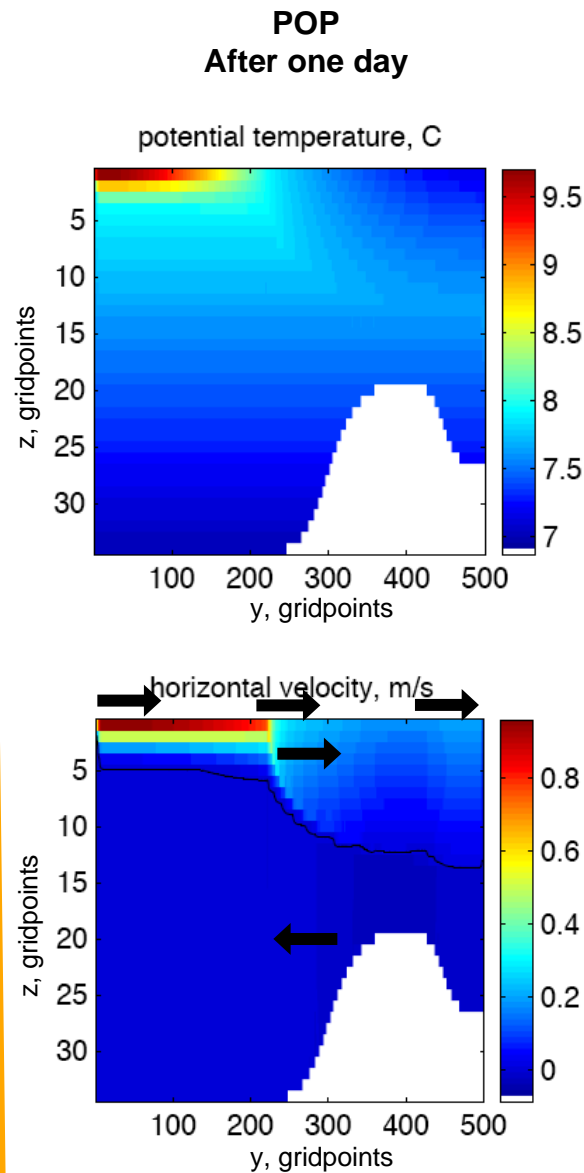
HyPOP: Reconciling barotropic and baroclinic SSH

- POP and HyPOP use barotropic/baroclinic splitting: implicit barotropic, explicit baroclinic, with the same timestep.
- Barotropic SSH includes surface gravity waves
- Layer thicknesses, which are treated as a tracer, do not include surface gravity waves.
- Sum of baroclinic layer thickness does not match the barotropic SSH.
- Correction: Could simply stretch $\sum_{k=1}^K h_k$ to match $H + \eta$, but this causes spurious diapycnal mass flux.
- We implemented a flux correction to the baroclinic layer thicknesses.



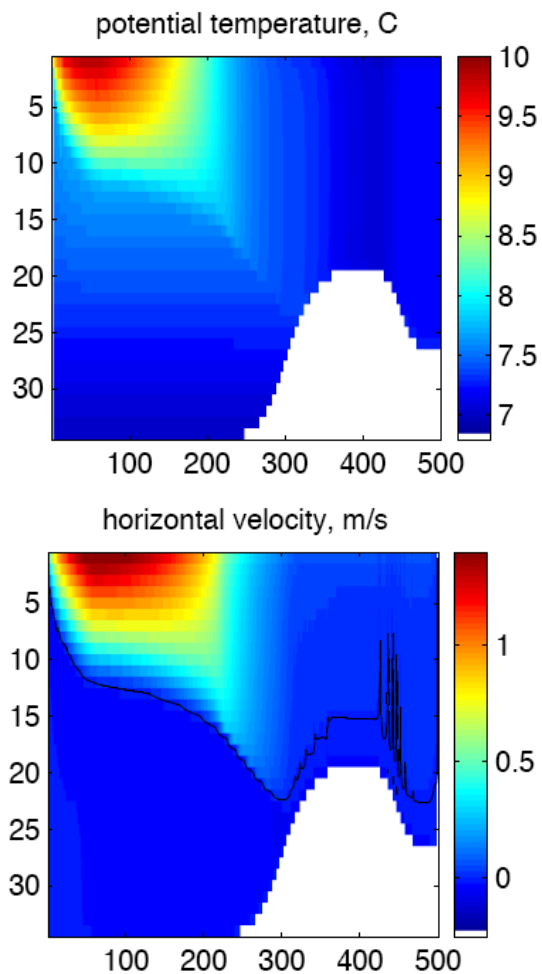
Overflow Domain and Forcing

- 2D domain, in y-z
- Coriolis force is off
- surface forcing of 12C to 2C
- surface wind stress
- vertical diffusion using kpp

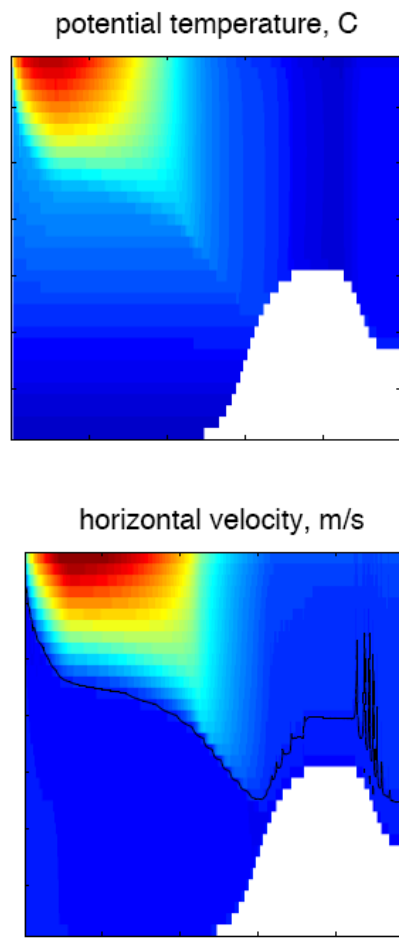


Simulation at Day 6

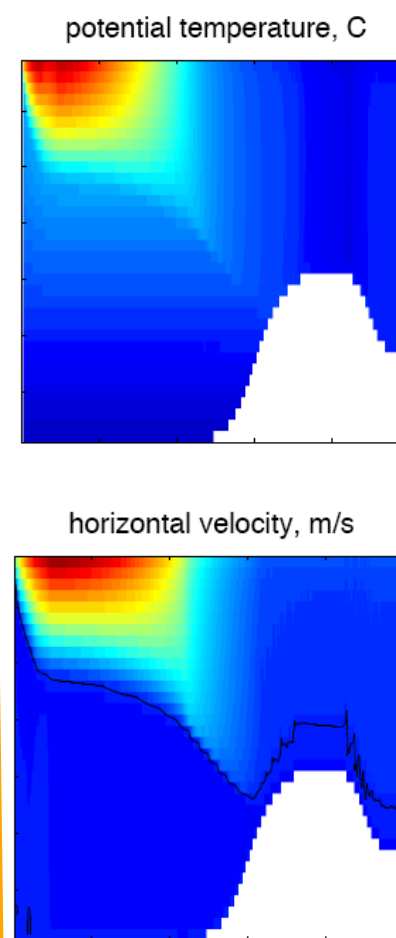
POP
One grid: Z-level
centered diff. advection



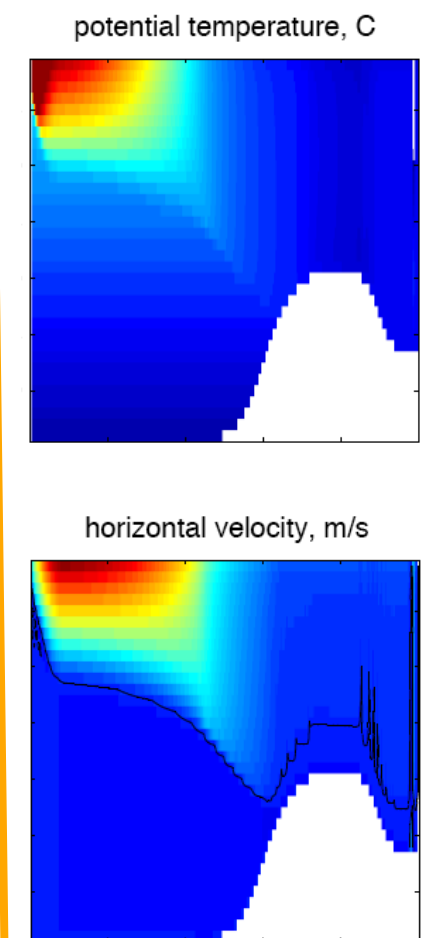
HyPOP in POP mode
momentum: Z-level
centered diff. advection
tracers: Z-level
centered diff. advection



HyPOP
momentum: Z-level
centered diff. advection
tracers: ALE, remap to Z
every 10 steps
incremental remap adv.



HyPOP
momentum: Z-level
centered diff. advection
tracers: ALE,
no remapping,
incremental remap advection



HyPOP: Conclusions

- **HyPOP infrastructure development is largely complete.**
- **Recent items:**
 - Vertical diffusion (KPP, Richardson number) conducted on union grid.
 - Cubic spline interpolation of density, integrate to compute pressure.
 - Reconciling SSH from the sum of baroclinic layer thickness with SSH from barotropic mode: use flux correction.
- **Next:**
 - Time-dependant regridding (relaxation to target grid)
 - Further testing with different vertical grids for momentum and tracers
- **Then moving into research and testing, including:**
 - Evaluate various grid combinations
 - Criteria for transition from Z-level in mixed layer to isopycnal grid at depth
 - Evaluation of HyPOP in global simulations
 - Improving efficiency of the code