Overview of Polynomial Chaos Methods for Uncertainty Quantification with Application to HYCOM

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Polynomial Chaos

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What is polynomial chaos? How does polynomial chaos work?

2 Uncertain inflow through Yucatan Straits.

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What is polynomial chaos?

- Idea of polynomial chaos originated with Norbert Wiener in 1938 — before computers.
- It is being used by engineers to assess how uncertainties in a model's inputs manifest in its outputs.
- It can be much more efficient than Monte Carlo methods.
- · Can it be useful to oceanographers?

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Why "polynomial"? Why "chaos"?

- "Chaos" simply refers to uncertainty. Nothing to do with strange attractors?
- Want to compute how uncertainties of a dynamical system's inputs manifest in its outputs.
- "Polynomial" refers to use of polynomial expansions to propagate uncertainties.
- Idea is to exploit orthogonality of the polynomials.

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- Initial conditions.
- Boundary conditions.
- Forcing.
- Parameters.
- Polynomial Chaos can handle only a limited number of uncertain inputs.
- But it focuses on all likely values of those few uncertain inputs.

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- Every field at every time.
- Value of a particular field at a particular point and a particular time.
- Derived quantities,
 e.g. maximum of the Meridional overturning stream function.
- Polynomial Chaos allows focus to be on points of interest.
- Not necessary to explore all uncertainties simultaneously.

- Call the uncertain input ξ and the output ϕ .
- Uncertainty of ξ is specified via its pdf $\rho(\xi)$.
- Want pdf of ϕ or at least information about how it varies as ξ varies.
- Basic Idea: Express output as a polynomial series.

$$\phi(\xi) = \phi_0 + \phi_1 P_1(\xi) + \phi_2 P_2(\xi) + \dots$$

• Orthogonal polynomials P_k are related to pdf ρ .

$$\int P_j(\xi)P_k(\xi)\rho(\xi)d\xi=\delta_{j,k}$$

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$$\phi(\xi) = \phi_0 + \phi_1 P_1(\xi) + \phi_2 P_2(\xi) + \dots$$
$$\int P_j(\xi) P_k(\xi) \rho(\xi) d\xi = \delta_{j,k}$$

- Is the series guaranteed to converge?
- In practice, it must be truncated.
- How to compute the coefficients $\phi_0, \phi_1, \phi_2, \dots$?

$$\phi_k = \frac{1}{N_k} \int \phi(\xi) P_k(\xi) \rho(\xi) d\xi$$

$$N_k = \int P_k^2(\xi) \rho(\xi) d\xi$$

Use Gaussian quadrature to evaluate the integrals.

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$$\phi_k = \frac{1}{N_k} \int \phi(\xi) P_k(\xi) \rho(\xi) d\xi$$

Gaussian Quadrature:

$$\int \phi(\xi) P_k(\xi) \rho(\xi) d\xi \approx \sum_p \phi(\xi_p) P_k(\xi_p) w_p$$

- Quadrature points: ξ_p
- Quadrature weights: w_p
- Computing outputs $\phi(\xi_p)$ for inputs at quadrature points ξ_p requires multiple model runs.
- How many quadrature points (runs) are needed?

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Two outputs:

$$\phi(\xi) = \phi_0 + \phi_1 P_1(\xi) + \phi_2 P_2(\xi) + \dots$$

$$\psi(\xi) = \psi_0 + \psi_1 P_1(\xi) + \psi_2 P_2(\xi) + \dots$$

- Two or more outputs require no more runs than does one output.
- Just save values for all outputs of interest, ϕ, ψ, \dots
- More coefficients, so more quadrature integrals are needed.
- Quadrature integrals are computationally cheap.
- Can examine uncertainty of an entire field.

$$\langle \phi \rangle = \int \phi(\xi) \rho(\xi) d\xi = \phi_0$$

variance:

$$\langle (\phi - \phi_0 \rangle)^2 = \sum_{k=1}^{k \text{max}} \phi_k^2$$

covariance:

$$\langle (\phi - \phi_0) (\psi - \psi_0) \rangle = \sum_{k=1}^{k_{\text{max}}} \phi_k \psi_k$$

Generate a cheap ensemble.

$$\phi(\xi) = \phi_0 + \phi_1 P_1(\xi) + \phi_2 P_2(\xi) + \dots$$

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- Call the uncertain inputs ξ and χ .
- Uncertainties of are specified via joint pdf $\rho(\xi,\chi)$.
- Now have polynomial series in two variables:

$$\phi(\xi,\chi) = \phi_0 + \phi_1 P_1(\xi,\chi) + \phi_2 P_2(\xi,\chi) + \dots$$

If uncertain inputs are independent, pdf factors:

$$\rho(\xi,\chi) = \rho_{\xi}(\xi)\rho_{\chi}(\chi)$$

- Then 2D quadrature reduces to two 1D quadratures.
- Number of quadrature points (runs) is squared.
- Curse of dimensionality.
- Sparse cubature might provide economy when exploring consequences of several uncertain inputs.

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Problem: How do uncertainties of the Yucatan inflow manifest within the Gulf of Mexico?

- Need to quantify inflow uncertainties.
- Inflow is characterized by several 2D time-varying fields.
- Computational cost increases dramatically with number of uncertain parameters.
- How to characterize uncertainties of inflow with only a few parameters?

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How to characterize uncertainties of inflow with only a few *independent* parameters?

- Use multivariate EOFs to characterize 2D spatial patterns of inflow uncertainty.
- Use corresponding principal components to characterize their temporal variability.
- Each mode's amplitude is assumed to have a Gaussian pdf.
- Hermite polynomials Gauss-Hermite quadrature.
- Quadrature points dictate the required HYCOM runs.
- Each run is the sum of a "favorite" inflow and its particular EOF contributions.

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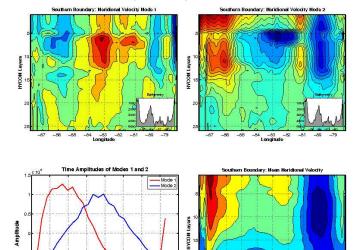
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Meridional velocity component of EOF.

100

200 250 300 350 Days



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now:

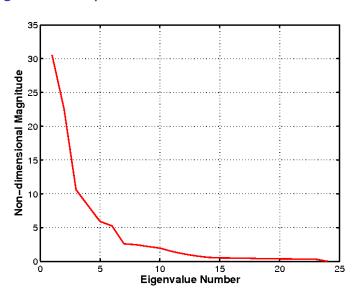
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- -83 -82 -81 -80 Longitude

-85 -85

Eigenvalue spectrum.



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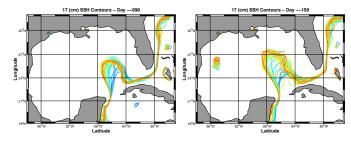
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Contours of 17 cm SSH from quadrature ensemble of 17 cm runs.



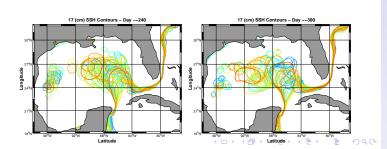


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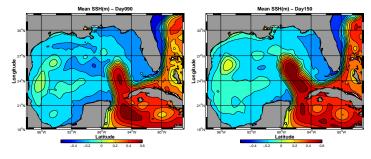
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Mean sea-surface height.



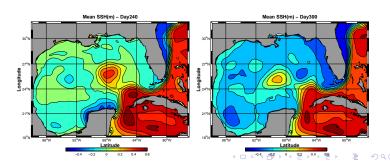


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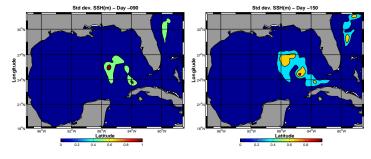
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Standard deviation of sea-surface height.





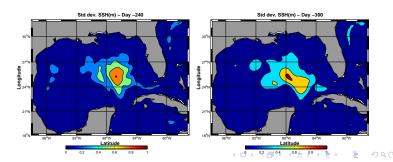
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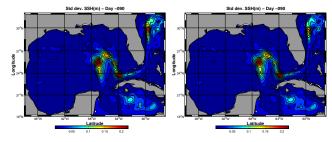
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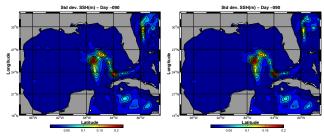
THOW example Setup.

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Convergence of series for SSH standard deviation for day 90.





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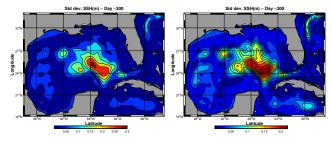
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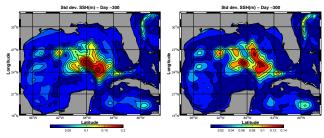
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Convergence of series for SSH standard deviation for day 300.





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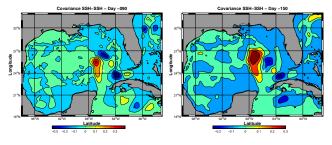
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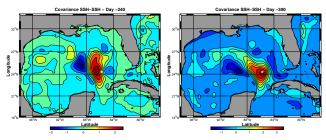
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Covariance of SSH with SSH at one point for days 90, 150, 240, and 300.





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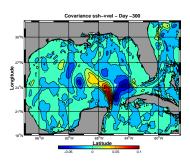
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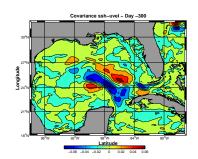
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Covariance of u,v-velocity with SSH at one point for day 90.





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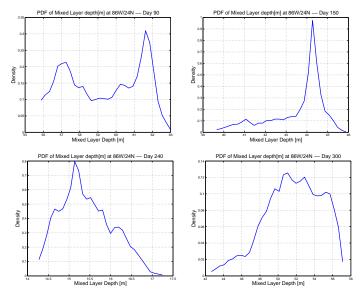
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Kernel density estimates for mixed-layer depth at day 90 from *artificial ensemble*.



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